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15.023J / 12.848J / ESD.128J Global Climate Change: Economics, Science, and Policy
Spring 2008

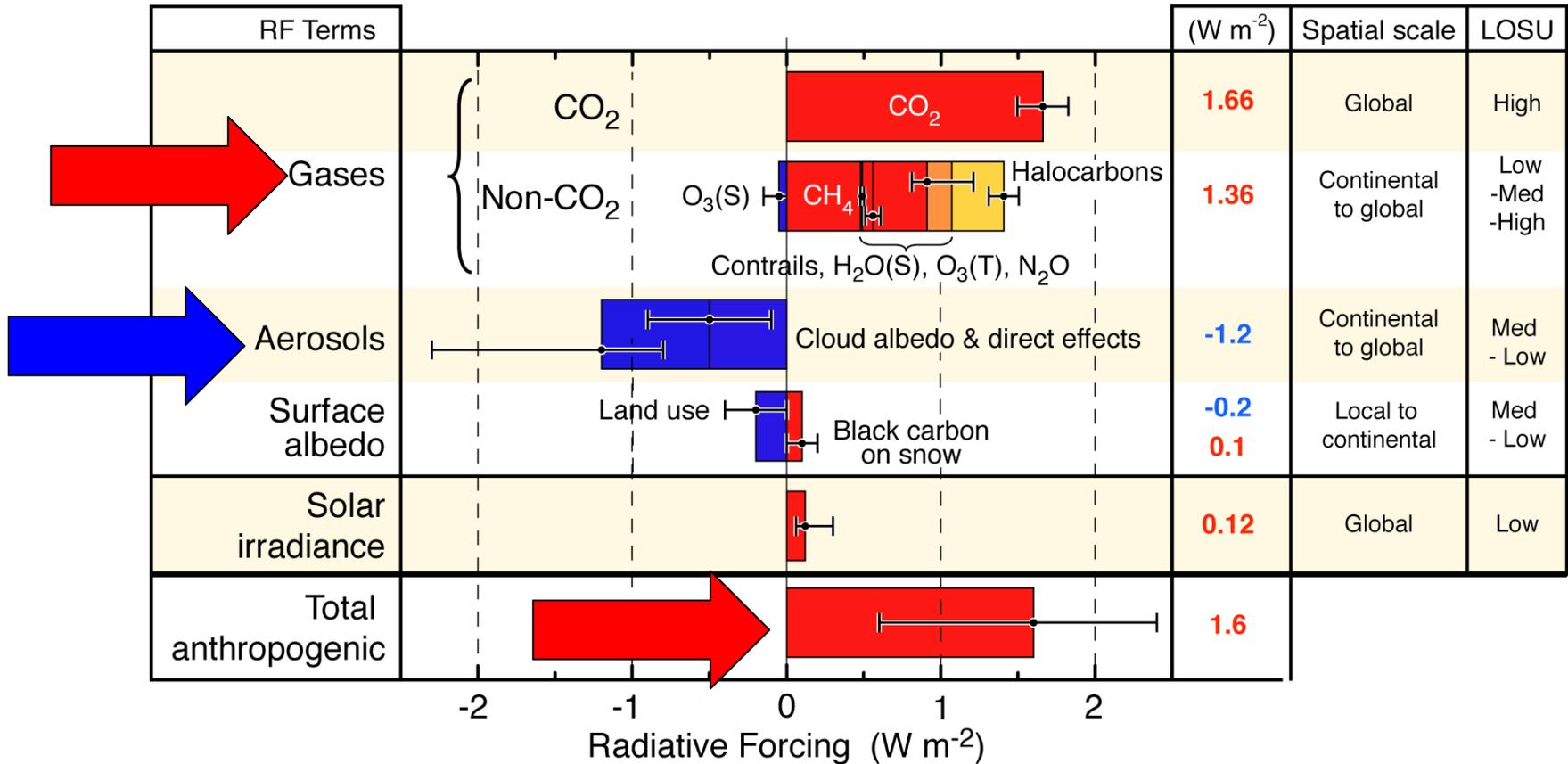
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THE CLIMATE MACHINE II :
Greenhouse Gas Exchange Rates and
Dynamics of the Atmosphere and Oceans

R. PRINN, February 25, 2008

- 1. Multiple greenhouse gases and their exchange rate**
- 2. Horizontal Circulations**
- 3. Vertical Circulations**
- 4. Ocean-atmosphere Coupling**
- 5. Fundamental Equations
& Model Integration**
- 6. Accuracy of Coupled Models**

WHAT ARE THE MAJOR HUMAN & NATURAL ACTIVITIES FORCING CLIMATE CHANGE IN THE INDUSTRIAL ERA (1750-2005)?



$1.6 \text{ W m}^{-2} \times 5.1 \times 10^{14} \text{ m}^2 = 8.16 \times 10^{14} \text{ W} = 816 \text{ TW}$ (about 52 times current global energy consumption)!

Ref: adapted from IPCC 4th Assessment, Summary for Policymakers, Feb. 2, 2007

THE KYOTO PROTOCOL REGULATES EMISSIONS OF CARBON DIOXIDE AND FIVE OTHER GREENHOUSE GAS CATEGORIES



Sources and Sinks of Regulated Gases Other Than CO₂

1. **Methane (CH₄)**
sources (biogenic, fossil) ←
sinks (OH)
2. **Nitrous oxide (N₂O)**
sources (biogenic, industrial) ←
sinks (ultraviolet radiation)
3. **Sulfur hexafluoride (SF₆)**
sources (industrial, natural) ←
sinks (extremely stable)
4. **Hydrofluorocarbons (HFCs) and
hydrochlorofluorocarbons (HCFCs)** ←
sources (industrial, natural)
sinks (OH)
5. **Perfluorocarbons (PFCs)** ←
sources (industrial, natural)
sinks (extremely stable)

HOW CAN WE
COMPARE
EMISSION
REDUCTIONS OF
NON-CO₂ GASES
TO CO₂ FOR
POLICY
PURPOSES?

THE KYOTO
PROTOCOL HAS
ADOPTED GLOBAL
WARMING
POTENTIALS TO
DEFINE THE
"EXCHANGE RATES"
BETWEEN GASES
FOR EMISSION
REDUCTION
PURPOSES



Global Warming Potentials (GWPs)

$$\text{GWP} = \frac{\int_0^T I_{\text{gas}} M_{\text{gas}} dt}{\int_0^T I_{\text{CO}_2} M_{\text{CO}_2} dt}$$

I_{gas} = instantaneous radiative forcing by gas at time t [depends on basic molecular properties and on atmospheric composition (gases, clouds, aerosols)]

M_{gas} = amount of added gas still remaining at time t [depends on lifetime of the gas, which in turn usually depends on amounts of the gas itself and of other gases (indirect effects)]

T = time horizon for integration [gases with lifetimes (longer/shorter) than CO₂ have GWPs (increasing/decreasing) with T]

Notes

1. Corrections must be made to account for effects of added gas on amounts of other greenhouse gases (e.g., effects of added CH₄ on O₃ and H₂O)
2. M_{CO_2} depends on assumptions regarding oceanic and terrestrial CO₂ sinks

GWP's VERY
 DEPENDENT
 ON THE GAS'S
 INSTANTANEOUS
 RADIATIVE
 FORCING, LIFETIME
 & CHOSEN
 TIME HORIZON
 KYOTO PROTOCOL
 ADOPTS A
 100-YEAR
 TIME
 HORIZON
 IS THIS
 SCIENTIFICALLY
 JUSTIFIABLE?
 (e.g. methane)

TABLE 3: Direct Global Warming Potentials (GWPs) relative to carbon dioxide (for gases for which the lifetimes have been adequately characterised). GWPs are an index for estimating relative global warming contribution due to atmospheric emission of a kg of a particular greenhouse gas compared to emission of a kg of carbon dioxide. GWPs calculated for different time horizons show the effects of atmospheric lifetimes of the different gases. [Based upon Table 6.7]

Gas	Lifetime (years)	Global Warming Potential (Time Horizon in years)		
		20 yrs	100 yrs	500 yrs
Carbon dioxide	CO ₂	1	1	1
Methane ^a	CH ₄	12.0 ^b	23	7
Nitrous oxide	N ₂ O	275	296	156
Hydrofluorocarbons				
HFC-23	CHF ₃	260	9400	12000
HFC-32	CH ₂ F ₂	5.0	1800	550
HFC-41	CHF ₂ F	2.6	330	97
HFC-125	CHF ₂ CF ₃	29	5900	3400
HFC-134	CHF ₂ CHF ₂	9.6	3200	1100
HFC-134a	CH ₂ FCF ₃	13.8	3300	1300
HFC-143	CHF ₂ CH ₂ F	3.4	1100	330
HFC-143a	CF ₃ CH ₃	52	5500	4300
HFC-152	CH ₂ FCH ₂ F	0.5	140	43
HFC-152a	CH ₃ CHF ₂	1.4	410	120
HFC-161	CH ₃ CH ₂ F	0.3	40	12
HFC-227ea	CF ₃ CHFCF ₃	33	5600	3500
HFC-236cb	CH ₂ FCF ₂ CF ₃	13.2	3300	1300
HFC-236ea	CHF ₂ CHFCF ₃	10	3600	1200
HFC-236fa	CF ₃ CH ₂ CF ₃	220	7500	9400
HFC-245ca	CH ₂ FCF ₂ CHF ₂	5.9	2100	640
HFC-245fa	CHF ₂ CH ₂ CF ₃	7.2	3000	950
HFC-365mfc	CF ₃ CH ₂ CF ₂ CH ₃	9.9	2600	890
HFC-43-10mee	CF ₃ CHFCF ₂ CF ₃	15	3700	1500
Fully fluorinated species				
SF ₆		3200	15100	22200
CF ₄		50000	3900	5700
C ₂ F ₆		10000	8000	11900
C ₃ F ₈		2600	5900	8600
C ₄ F ₁₀		2600	5900	8600
c-C ₄ F ₈		3200	6800	10000
C ₅ F ₁₂		4100	6000	8900
C ₆ F ₁₄		3200	6100	9000
Ethers and Halogenated Ethers				
CH ₃ OCH ₃		0.015	1	1
HFE-125	CF ₃ OCHF ₂	150	12900	14900
HFE-134	CHF ₂ OCHF ₂	26.2	10500	6100
HFE-143a	CH ₃ OCHF ₃	4.4	2500	750
HCFE-235da2	CF ₃ CHClOCHF ₂	2.6	1100	340
HFE-245fa2	CF ₃ CH ₂ OCHF ₂	4.4	1900	570
HFE-254cb2	CHF ₂ CF ₂ OCH ₃	0.22	99	30
HFE-7100	C ₄ F ₉ OCH ₃	5.0	1300	390
HFE-7200	C ₄ F ₉ OC ₂ H ₅	0.77	190	55
H-Galden 1040x	CHF ₂ OCF ₂ OC ₂ F ₄ OCHF ₂	6.3	5900	1800
HG-10	CHF ₂ OCF ₂ OCHF ₂	12.1	7500	2700
HG-01	CHF ₂ OCF ₂ CF ₂ OCHF ₂	6.2	4700	1500

^a The methane GWPs include an indirect contribution from stratospheric H₂O and tropospheric CH₄ production.
^b The values for methane and nitrous oxide are adjustment times, incorporating indirect effects of emission of each gas on its own lifetime.

***THE ATMOSPHERE AND OCEAN ARE
CIRCULATING FLUIDS WHICH CAN
TRANSPORT HEAT, MOMENTUM, MASS,
GREENHOUSE GASES AND PARTICLES
(cloud droplets, ice crystals, aerosols)
OVER GREAT DISTANCES.***

**THE DYNAMICS OF THESE FLUIDS
AND THE EXCHANGES BETWEEN THEM
PLAY A MAJOR ROLE IN CLIMATE.**

WHAT ARE THE HORIZONTAL CIRCULATIONS?

ATMOSPHERE

(lines are height (m)
of 1 bar surface)

Images removed due to copyright restrictions. See Figure 7.1a and Figure 8.1 in:

Peixoto, Jose, and Abraham Oort. *Physics of Climate*. New York, NY: Springer, 1992. ISBN: 9780883187128.

OCEAN

WHAT ARE THE VERTICAL CIRCULATIONS?

ATMOSPHERE

OCEAN

Images removed due to copyright restrictions. See Figure 10.2 in:

Peixoto, Jose, and Abraham Oort. *Physics of Climate*. New York, NY: Springer, 1992. ISBN: 9780883187128.

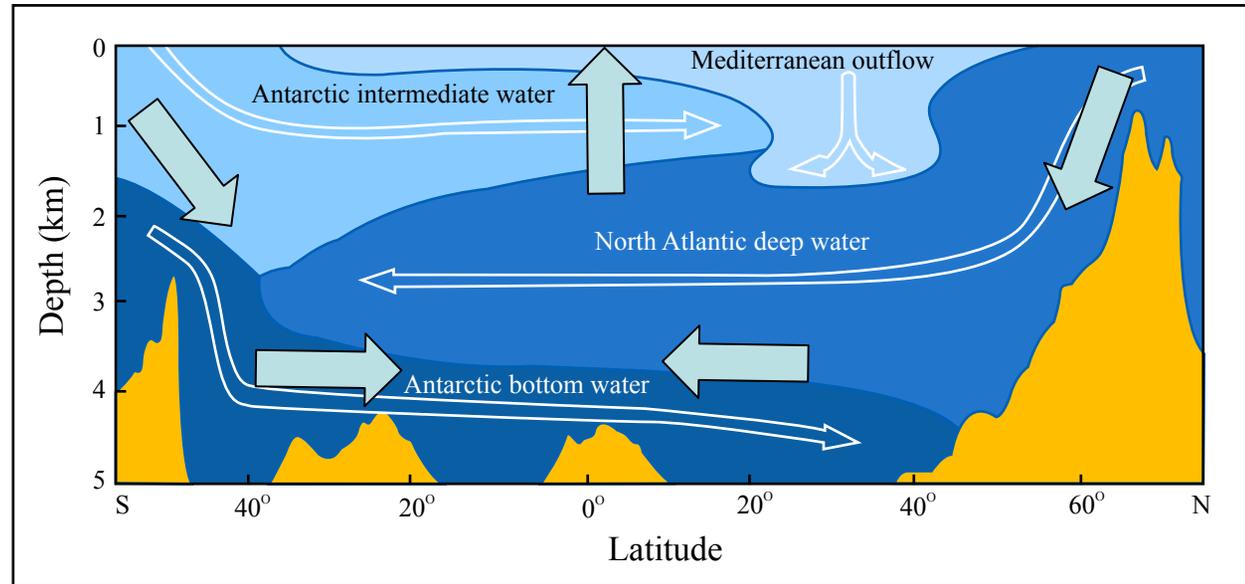
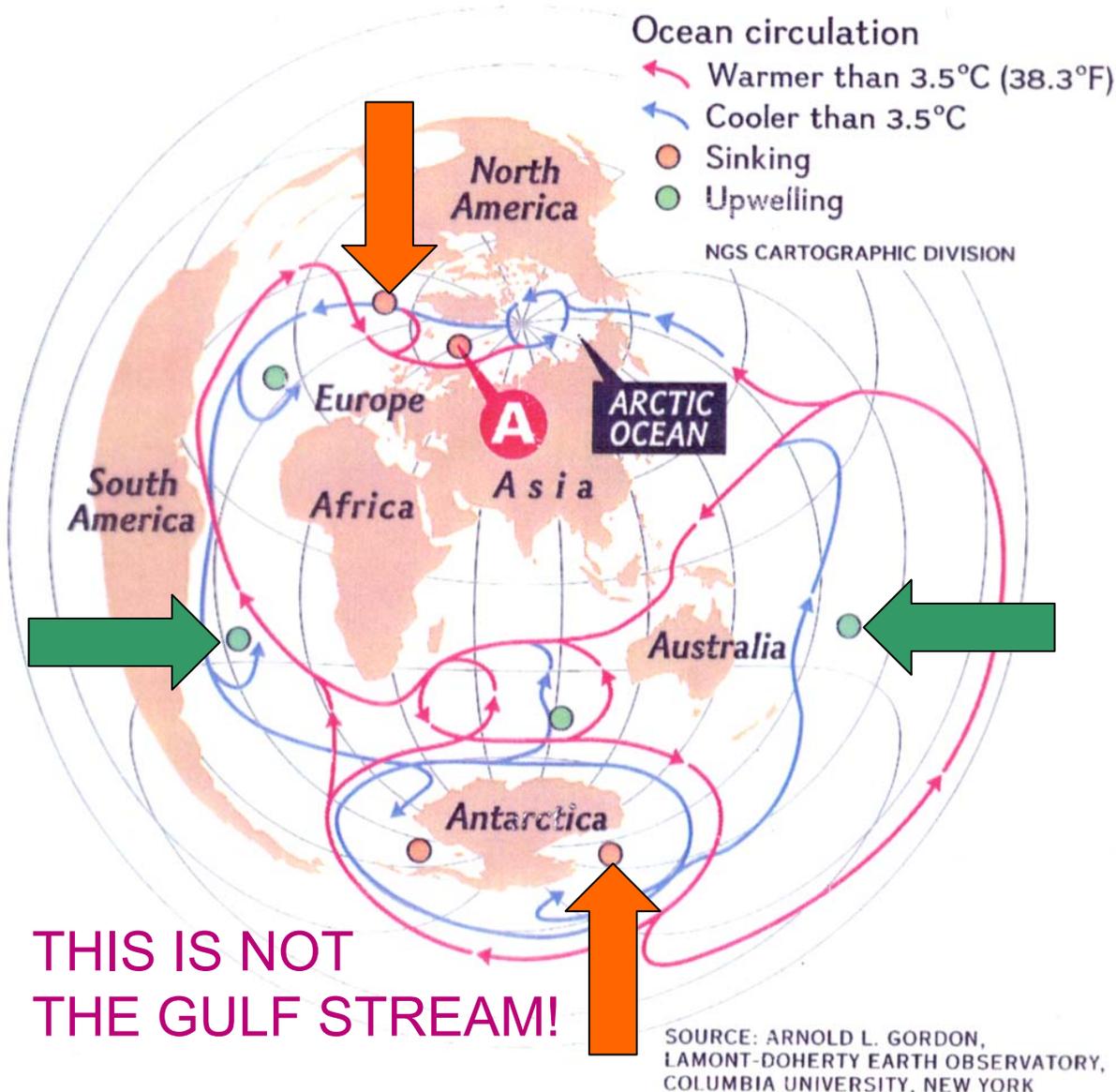


Figure by MIT OpenCourseWare.

Deep Atlantic circulation; water filling the North Atlantic basin comes from sources in the high-latitude North Atlantic, the Southern Ocean near Antarctica, and (at shallower depths) the Mediterranean Sea. (Adapted from E. Berner and R. Berner, *Global Environment* [Englewood Cliffs, N.J. Prentice -Hall, 1996].)

DEEP OCEAN OVERTURNING (THERMOHALINE) CIRCULATION



VERY IMPORTANT AS A HEAT AND CARBON DIOXIDE SINK

COMPRISES SINKING WATER IN THE POLAR SEAS (Norwegian, Greenland, Labrador, Weddell, Ross) AND RISING WATER ELSEWHERE

SLOWED BY DECREASED SEA ICE & INCREASED FRESH WATER INPUTS INTO THESE SEAS

INCREASED RAINFALL, SNOWFALL & RIVER FLOWS, & DECREASED SEA ICE, EXPECTED WITH GLOBAL WARMING!

RADIANT HEATING OR COOLING AND INFERRED HORIZONTAL HEAT TRANSPORT

AS A RESULT
OF THESE
CIRCULATIONS,
WHERE DOES
RADIANT
HEAT ENTER
AND LEAVE
THE EARTH?

USE SATELLITES TO
MEASURE DIFFERENCE
(IN WATT/M²)
BETWEEN INCOMING
SOLAR & OUTGOING
INFRARED RADIATION

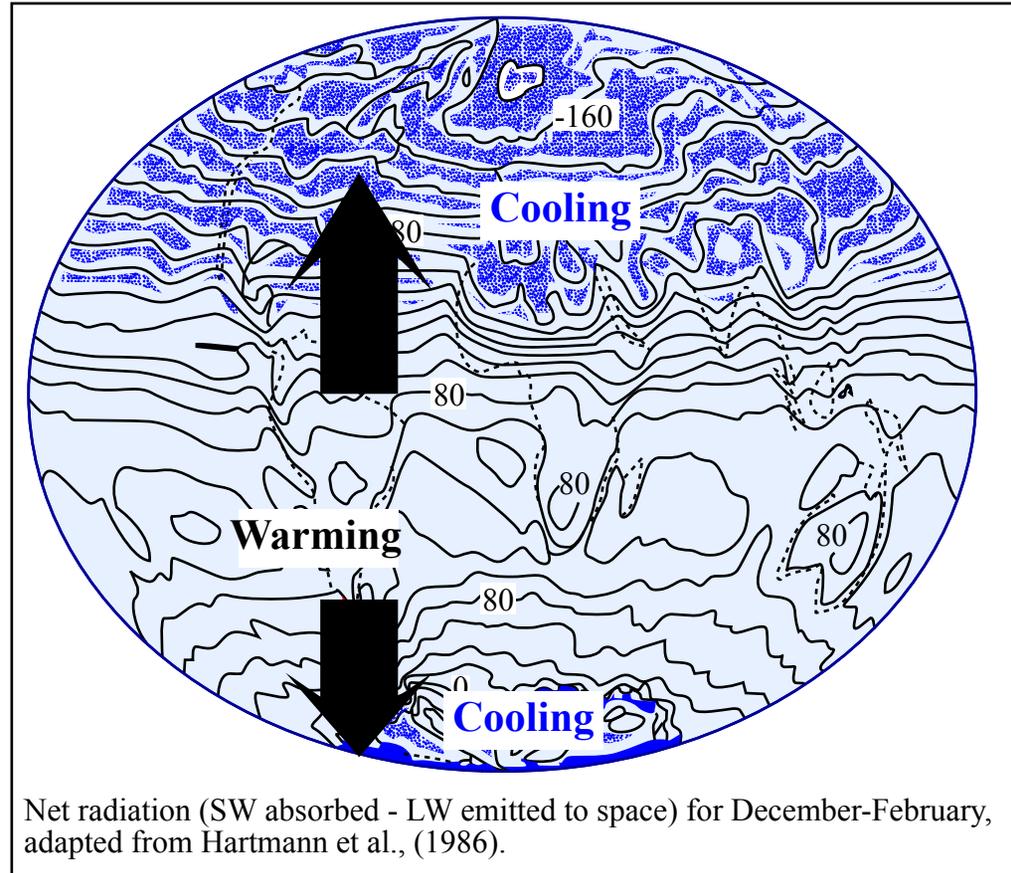


Figure by MIT OpenCourseWare.

Fundamental Equations

Models of the earth's climate are based on laws of physics:

- Conservation of energy (sensible (temp.), latent (evap./cond.), radiation)
- Conservation of momentum (north-south, east-west, up-down)
- Conservation of mass
- Equation of state (relates pressure, density and temperature)
- Conservation of chemical elements

Express the changes in the variables by “budget” or “continuity” equations:

$$\left(\frac{\partial \Phi}{\partial t} \right) = \frac{d\Phi}{dt} - \frac{\partial}{\partial x} (u\Phi) - \frac{\partial}{\partial y} (v\Phi) - \frac{\partial}{\partial z} (w\Phi)$$

The local change
(rate of accumulation)
of () in the box

Actual production
or destruction of ()
within the box

Change in () due to loss to downstream
boxes or arrival of () from an upstream
box (called advection or convection)

COUPLING OF GENERAL CIRCULATION MODELS (GCM's) OF THE OCEAN & ATMOSPHERE

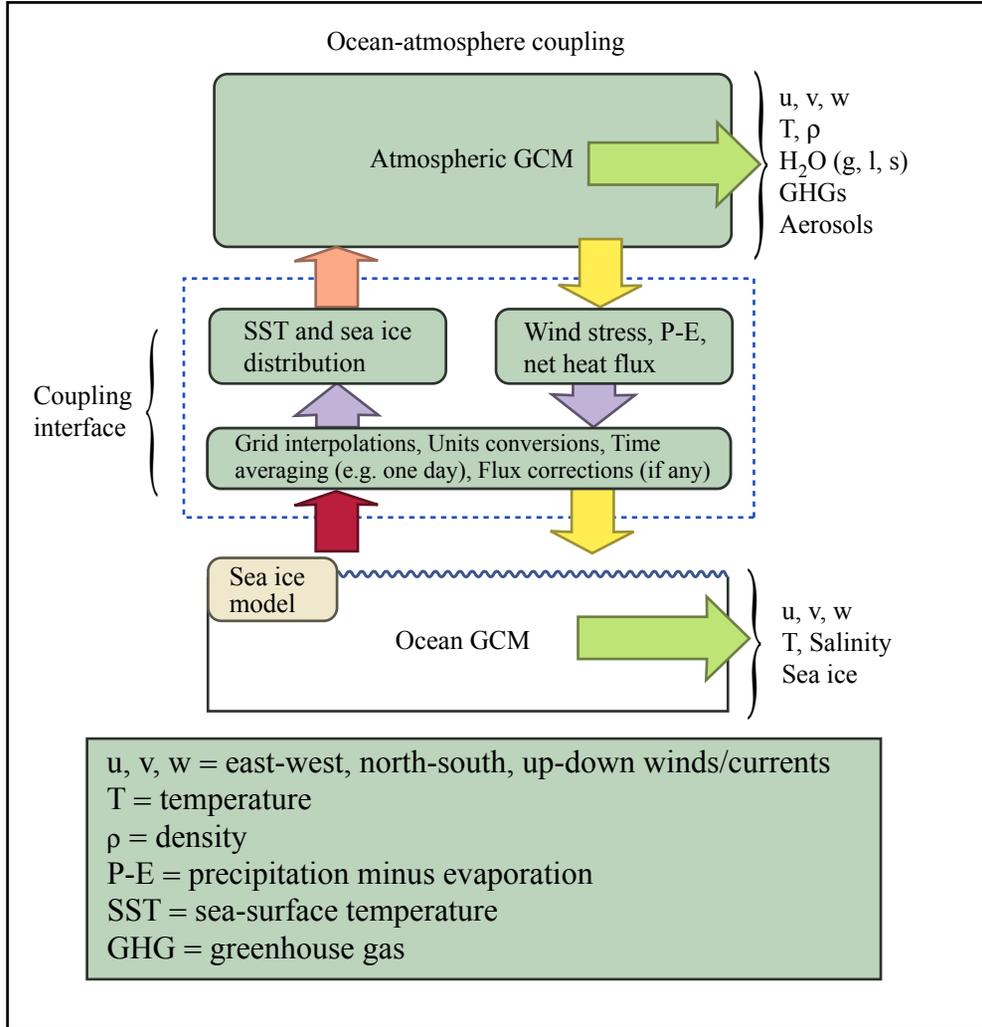
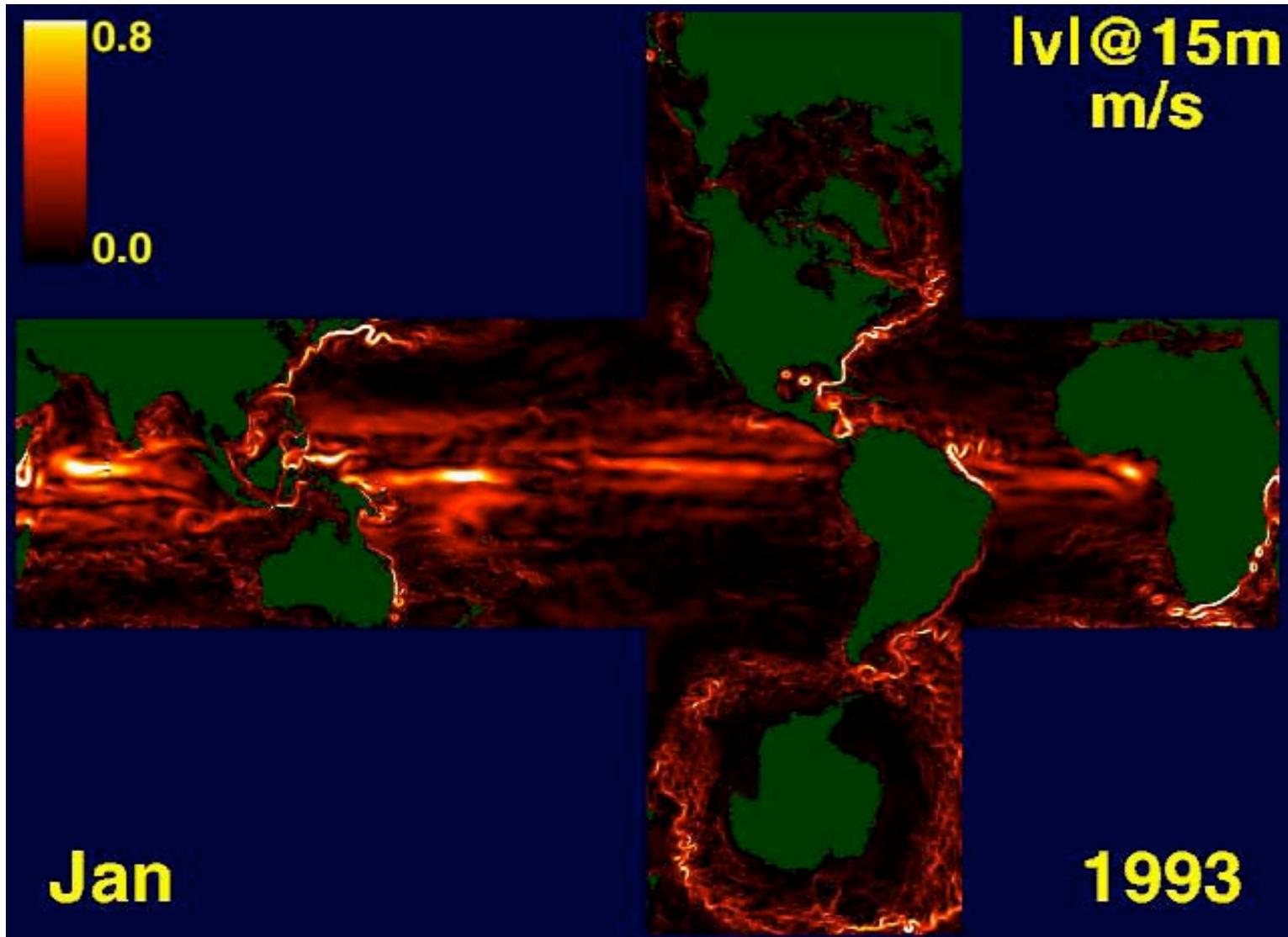


Figure by MIT OpenCourseWare.

AN EXAMPLE: OCEAN CIRCULATION MODEL (CGCS CLIMATE MODELLING INITIATIVE)

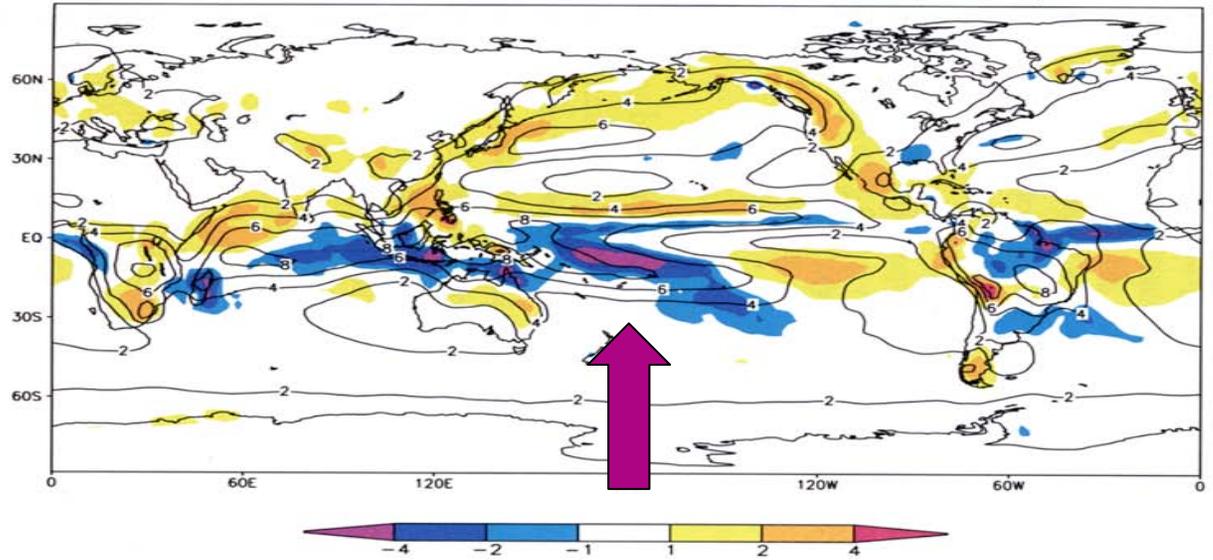


USES NOVEL CUBED SPHERE GRID

CURRENTS AT 15 METERS DEPTH

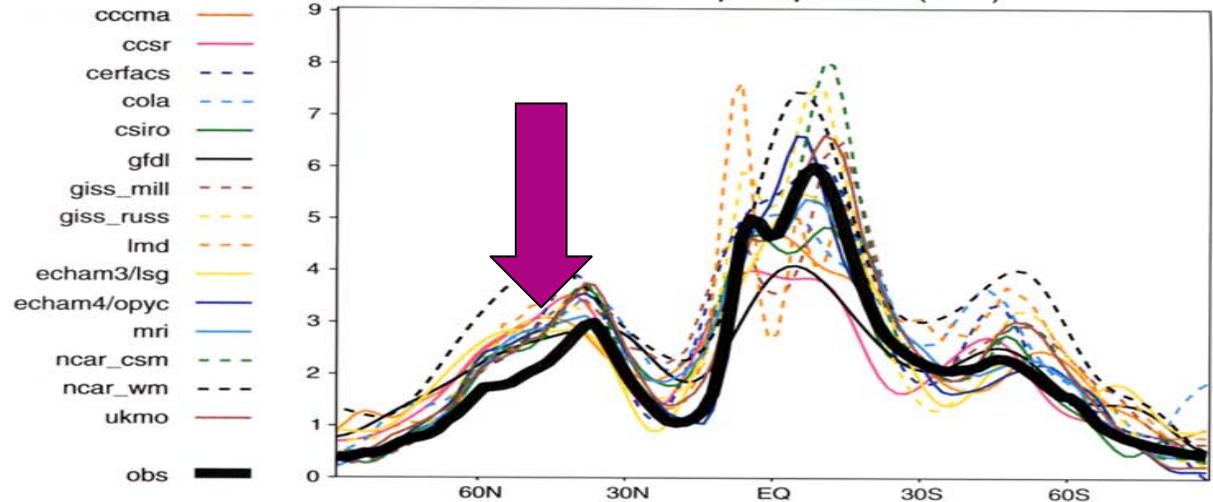
PROBLEMS WITH CONVECTION & CLOUDS

Model precipitation (DJF)
Model mean (contoured) mean minus observed (shaded)



ERRORS CAN EXCEED 4MM/DAY (1.5M/YEAR)!

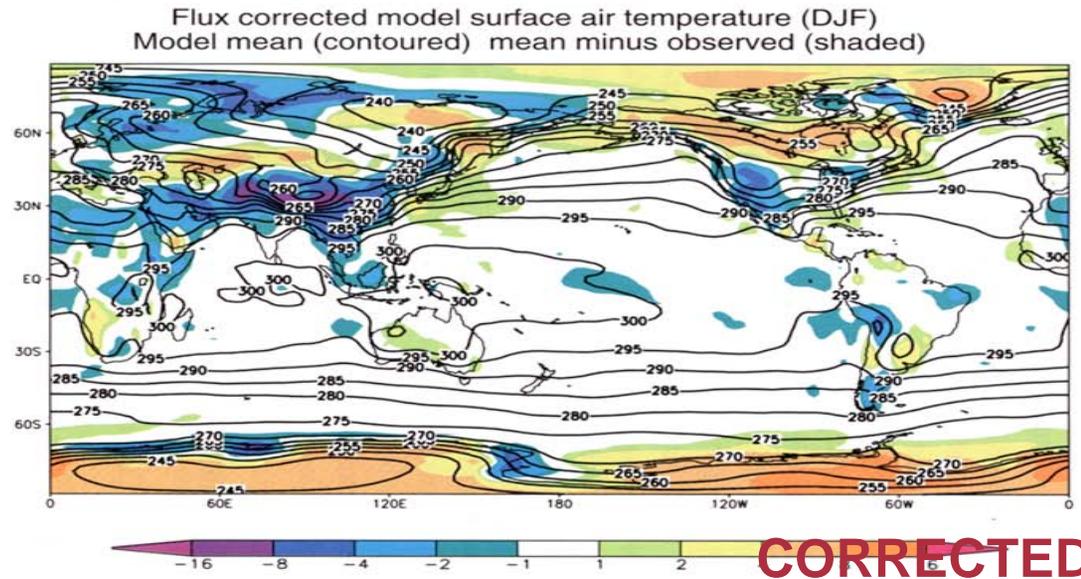
Zonal mean precipitation (DJF)



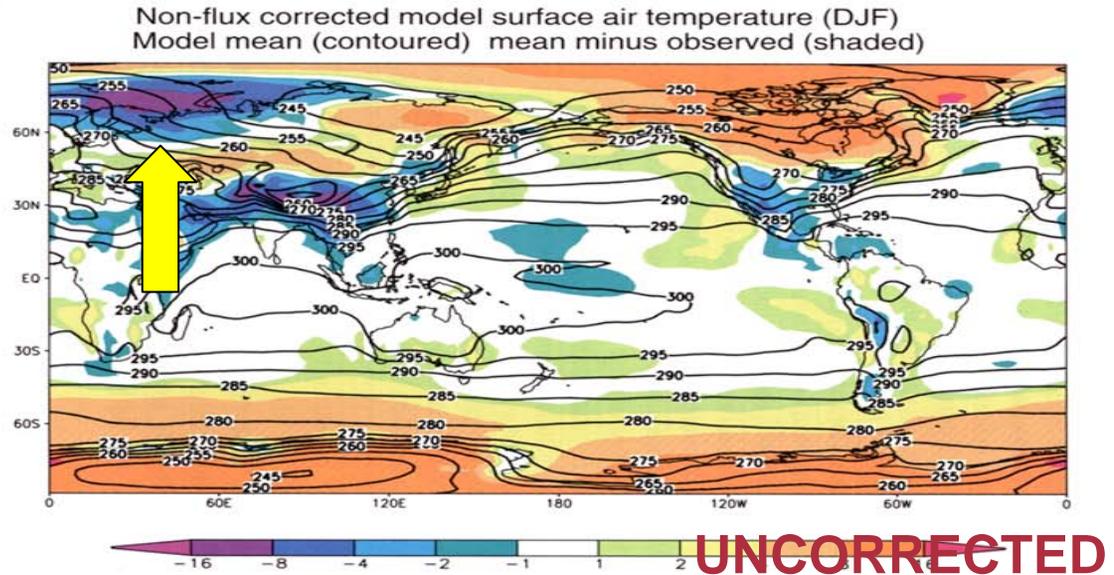
Courtesy of the Intergovernmental Panel on Climate Change. Used with permission.

(precipitation in mm/day)

PROBLEMS IN COUPLING (Flux adjustments)



WITHOUT THESE
ADJUSTMENTS
CAN HAVE
REGIONAL
TEMPERATURE
ERRORS
EXCEEDING 16°C!



Courtesy of the Intergovernmental Panel on Climate Change. Used with permission.

(temperatures in degrees Kelvin)

HAVE MODELS IMPROVED OVER THE PAST SIX YEARS?

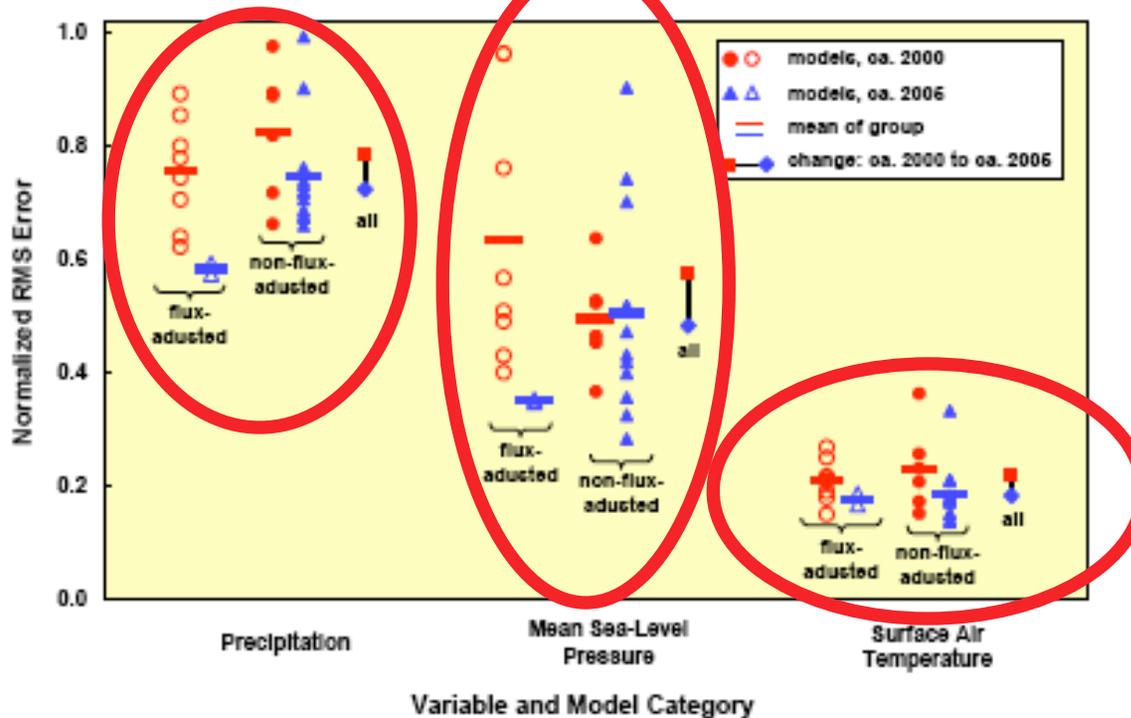


Figure 8.11. Normalized root-mean-square error in simulation of climatological patterns of monthly precipitation, mean sea-level pressure, and surface air temperature. Recent AOGCMs (ca. 2005) are compared to their predecessors (ca. 2000, and earlier). Models are categorized based on whether or not any flux adjustments were applied. The models are gauged against the following observation-based datasets: CMAP (Xie and Arkin, 1997) for precipitation (years 1980–1999), ERA40 (Uppala et al., 2005) for sea-level pressure (years 1980–1999), and CRU (Jones et al., 1999) for surface temperature (years 1961–1990). Before computing the errors, both the observed and simulated fields were mapped to a uniform 4 x 5 degree latitude-longitude grid. For the earlier generation of models, results are based on the archived output from control runs (specifically, the first 30 years, in the case of temperature, and the first 20 years for the other fields), and for the recent generation models, results are based on the 20th Century simulations with climatological periods selected to correspond with observations. (In both groups of models, results are insensitive to the period selected.)

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