

Biogenic opal

What is it?

Amorphous silica:



Precipitated in the surface ocean by:

- phytoplankton
 diatoms, silicoflagellates
- protozoans
 radiolaria

A fraction fall of this opal falls to the sea floor

-- it's efficiently recycled, in water column and sediments

-- overall, ~ 3% of opal production is preserved in sediments

The solubility of biogenic opal in seawater Initial Studies -- Hurd, 1973, GCA 37, 2257-2282

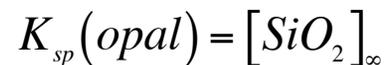
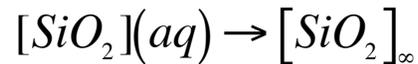
Experiment:

- separate opal from cores
- clean with acid
- place in sw at controlled temp

After ~ days:

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Please see: Hurd, D. "Interactions of biogenic opal, sediment, and seawater in the Central Pacific." *Geochimica et Cosmochimica Acta* 37 (1973): 2257-2282.



Note:

T dependence

Solubility ~ 900 μ M at 2°C !

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A mineral,
undersaturated in seawater
apparently simple dissolution kinetics...

What do we expect $[\text{Si}(\text{OH})_4]$ in pore water to look like?

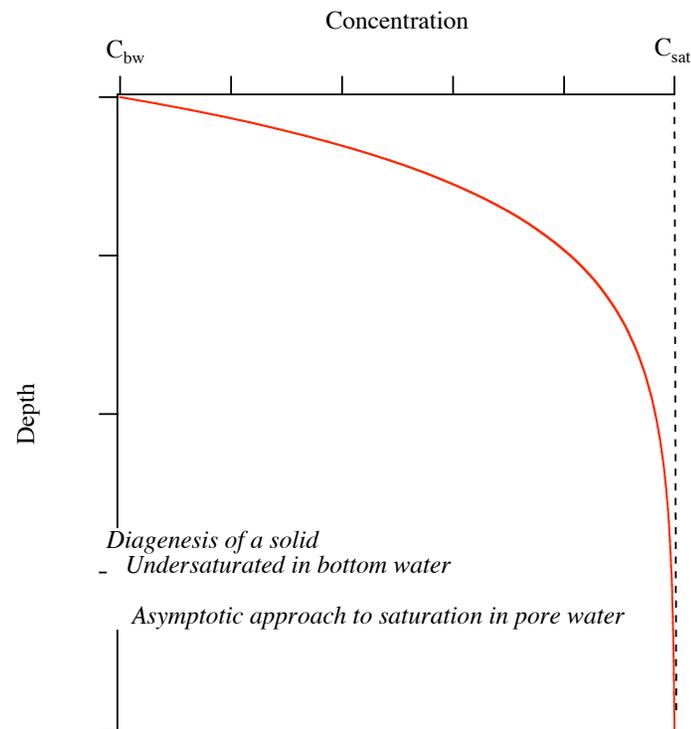
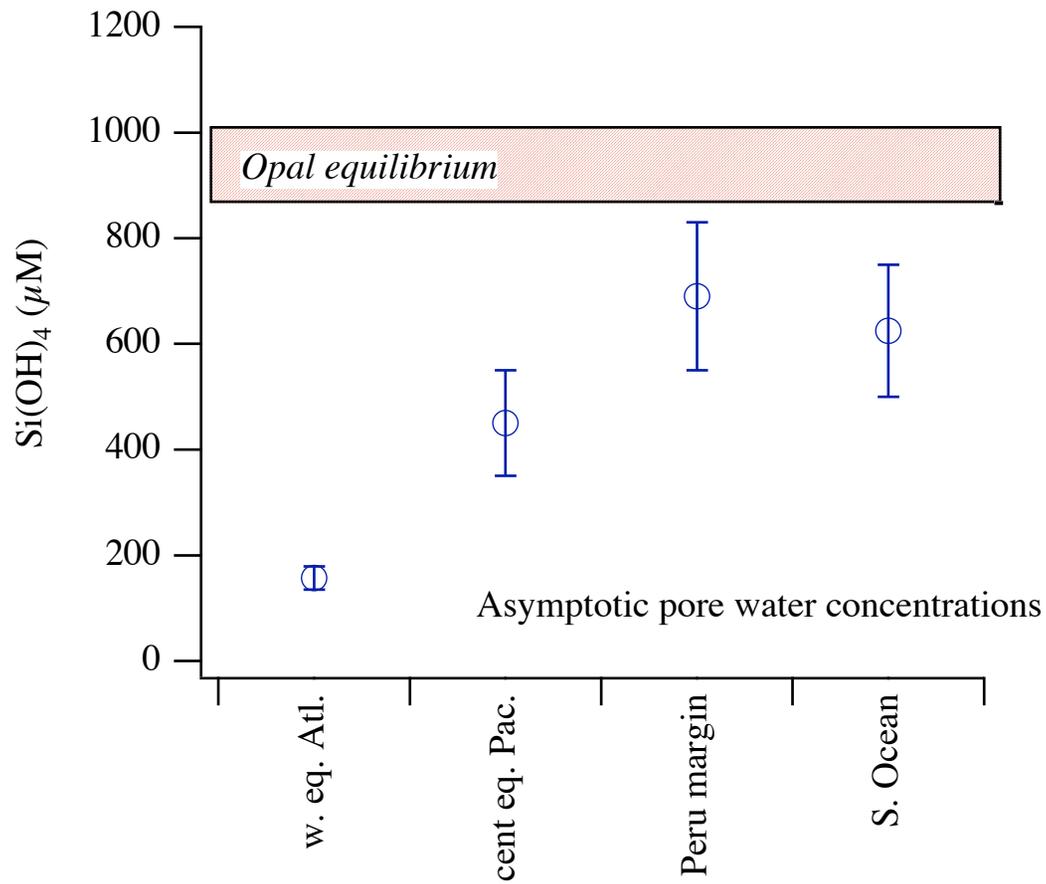


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Comparing asymptotic pore water [Si(OH)₄] to the “equilibrium” value



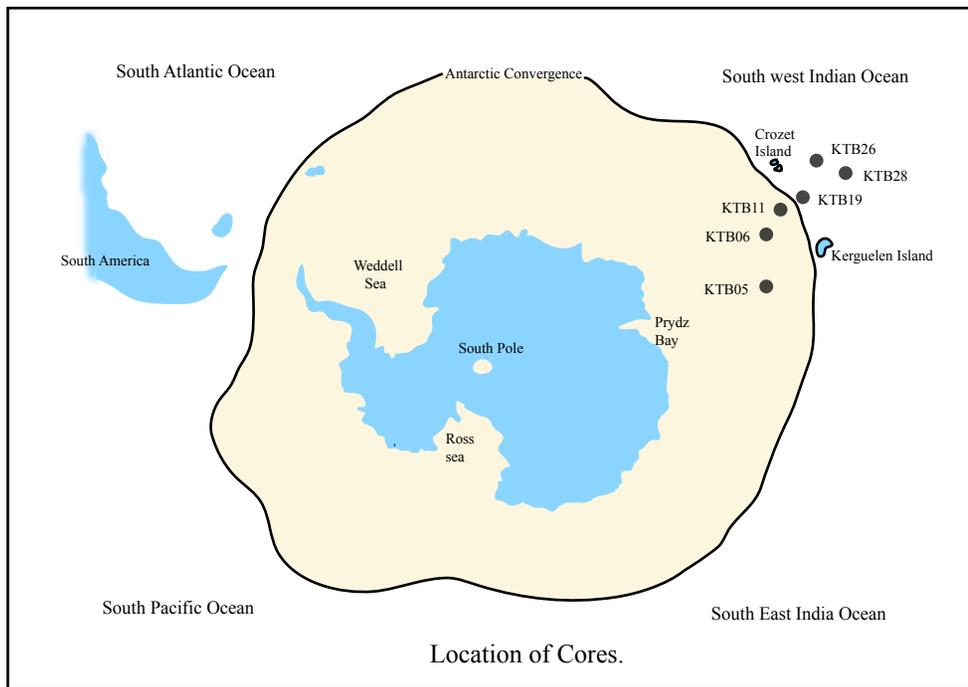


Figure by MIT OCW.

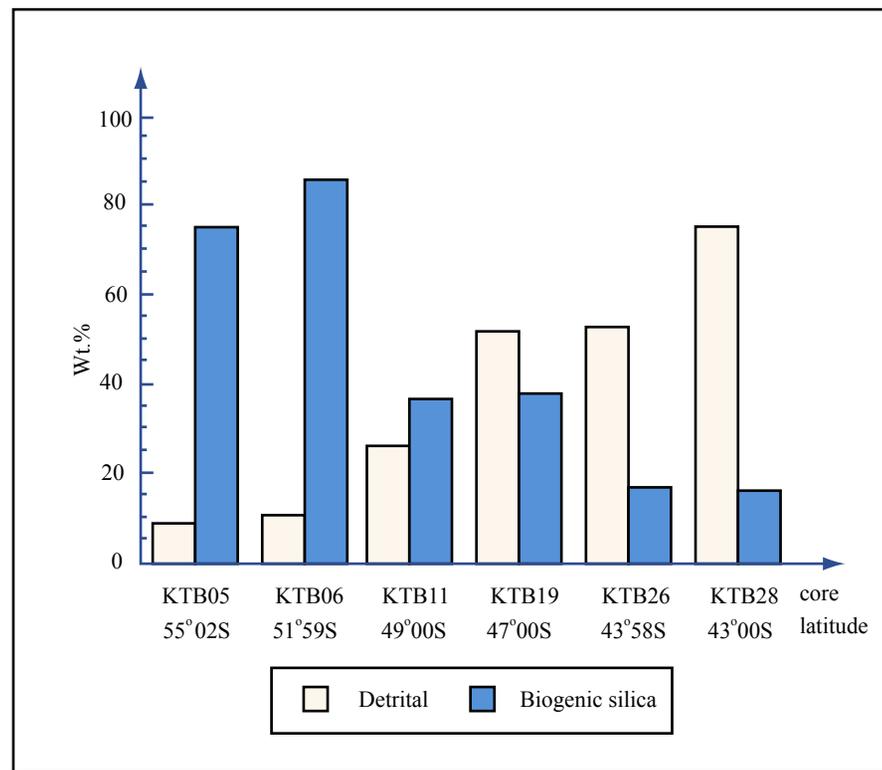


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Please see: Cappellen, Philippe Van, and Linqing Qiu. "Biogenic silica dissolution in sediments of the Southern Ocean." *J Solubility*, 1109-1128.

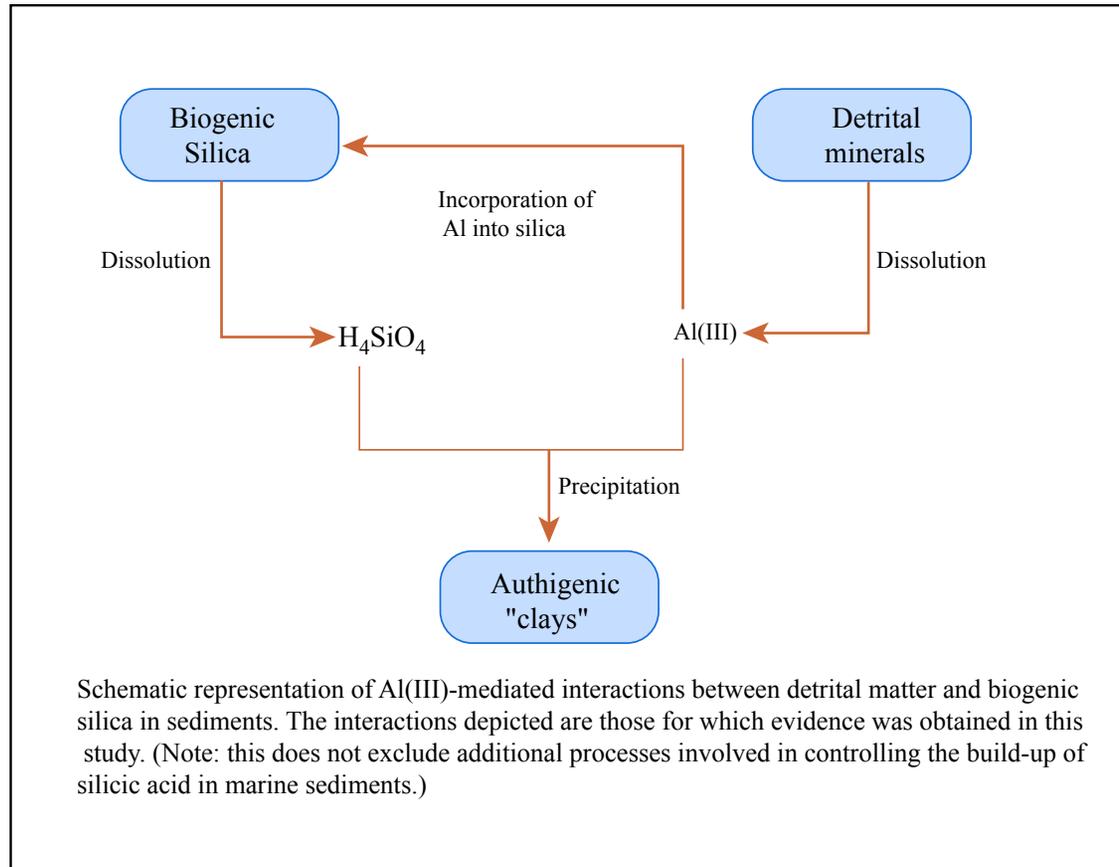


Figure by MIT OCW.

Studies of the Preservation Rate of Opal in deep-sea sediments

Components of the studies:

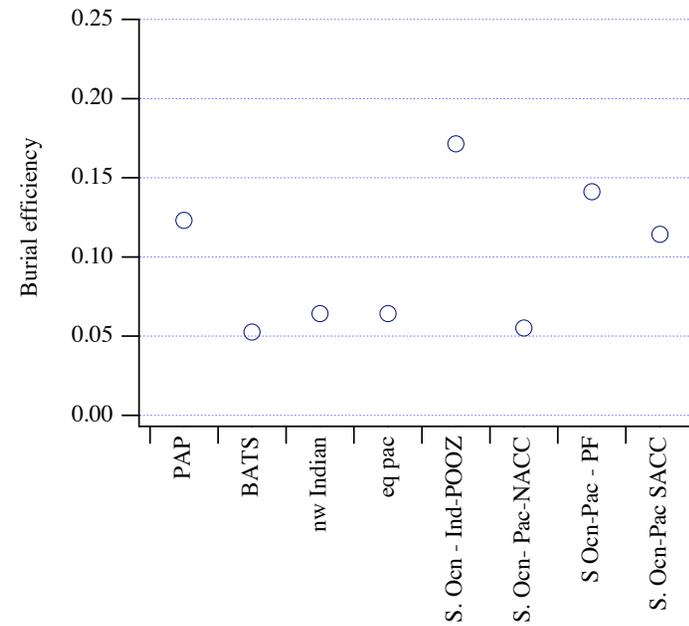
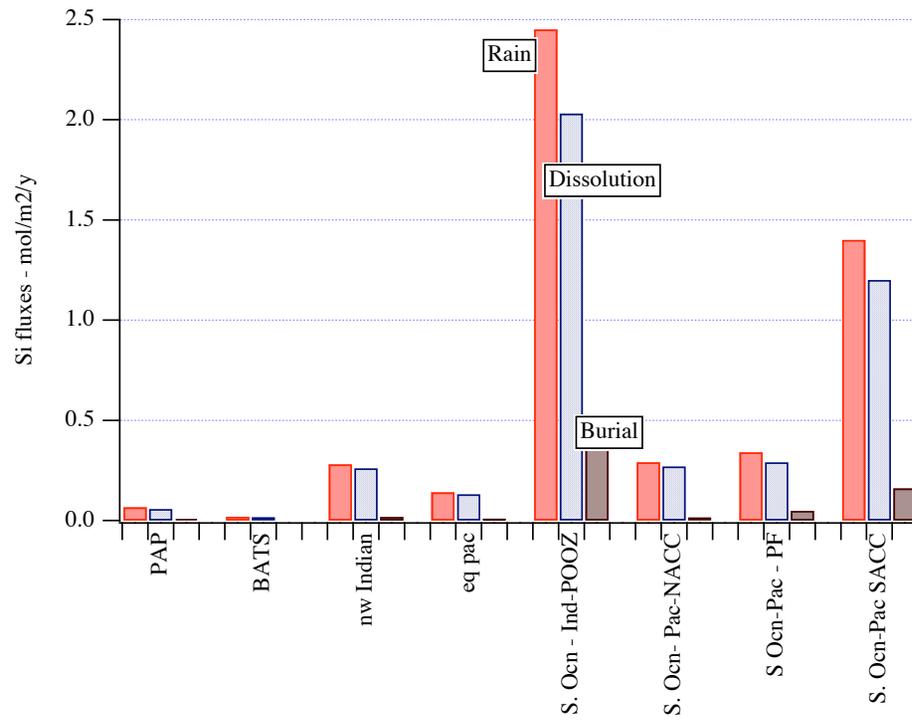
Rain rate to sea floor : time-series sediment traps

Benthic remineralization rates : flux chambers ; pore waters

Burial rates : solid phase measurements

Opal preservation efficiency in sediments : summary

Sources: Ragueneau et al., 2001
 Nelson et al., 2002, DSR II 49, 1645-1674



% opal in sediments -- from Gruber and Sarmiento

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Please see: Gruber, and Sarmiento. *Ocean Biogeochemical Dynamics*. Princeton, NJ: Princeton University Press, 2006. ISBN: 9780691017075.

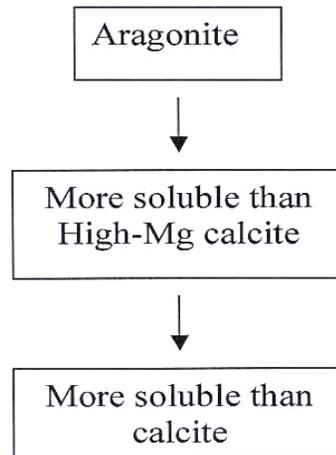
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Mechanism for CaCO₃ Dissolution in Sediments:

Dissolution in pore waters that are undersaturated w.r.t. the abundant mineral phase

1. Which mineral?



..... CALCITE is the predominant mineral in the deep sea.

2. Quantifying solubility in seawater

Conditional Solubility product,

$$K_{sp} = [Ca^{2+}][CO_3^{2-}]$$

T=25°C, P=1 atm, S=35 psu:

Ingle (1975) $4.60 \pm 0.10 \times 10^{-7}$

Mucci (1983) $4.30 \pm 0.20 \times 10^{-7}$

i.e., $\pm \sim 5\%$

The pore water based estimate (Sayles 1985)

- ① Measure $\text{Alk} + \text{Ca} \Sigma \text{CO}_2$, vs. depth in pore waters of sediments underlying water columns of varying depth
- ② Calculate $[\text{CO}_3^{2-}] \times [\text{Ca}^{2+}]$
: Most often is constant below a few cm.
- ③ If the pore waters have reached saturation with respect to calcite, $([\text{CO}_3^{2-}][\text{Ca}^{2+}])_{\text{meas}} = K_{sp}^{\text{sw}}$

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Please see: Sayles, J. "CaCO₃ solubility in marine sediments: Evidence for equilibrium and non-equilibrium behavior." *Geochim Cosmochim Acta* 49 (1985): 877-888.

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Temperature dependence of calcite solubility: Experimentally determined:

As T decreases, K_{sp} increases

Pressure dependence of calcite solubility: Both experimentally determined and calculated:

As P increases, K_{sp} increases

SO: CALCITE BECOMES MORE SOLUBLE WITH INCREASING DEPTH IN THE OCEAN.

Describing the saturation state of seawater & pore water with respect to calcite:

“Degree of saturation”

$$\Omega = \frac{([Ca^{2-}][CO_3^{2-}])_{meas}}{K_{sp}} \approx \frac{[CO_3^{2-}]_{meas}}{[CO_3^{2-}]_{equil}}$$

“Delta-carbonate”

$$\Delta CO_3^{2-} = [CO_3^{2-}]_{meas} - [CO_3^{2-}]_{equil} \quad \text{in } \mu\text{mol/kg}$$

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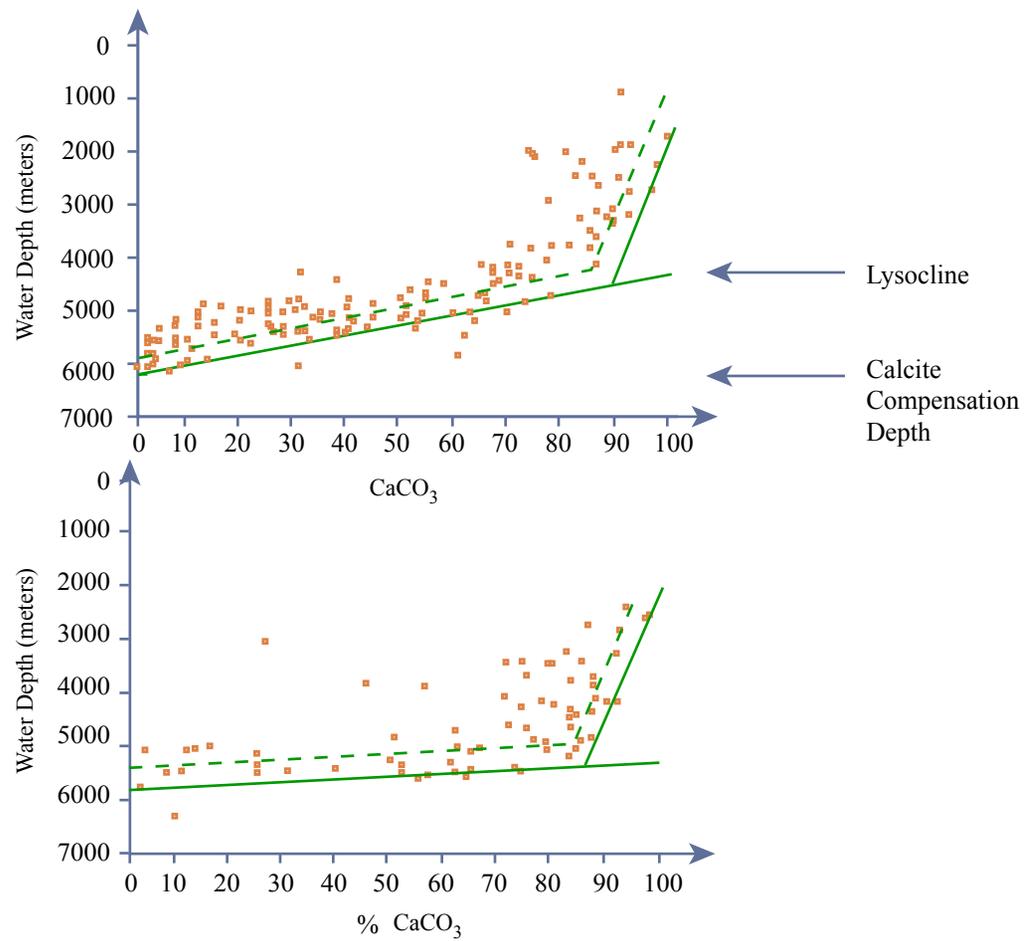
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Lysocline : depth at which there is evidence of dissolution

CCD : depth at which % CaCO₃ = 0 i.e, at which dissolution rate = supply rate.

CCD is straightforward; but what does lysocline indicate?

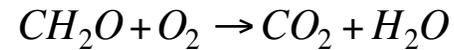
Is %CaCO₃ a sensitive indicator of dissolution?

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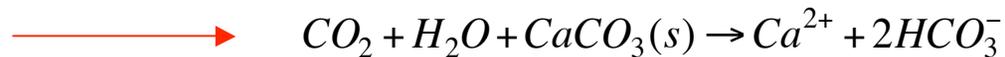
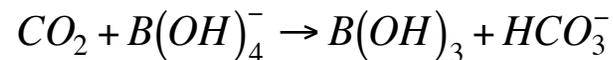
Please see: Broecker, W. S., and T.-H. Peng. *Tracers in the Sea*. Palisades, NY: Lamont-Doherty Geological Observatory, 1982.

“Metabolic” calcite dissolution

Oxic respiration results in the release of acids to solution :



Acids are neutralized by



(and similar reactions for neutralizing H⁺)

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Please see: Hales, B., et al. "Respiration and dissolution in the sediments of the western North Atlantic: Estimates from models of in situ microelectrode measurements of porewater oxygen and pH." *DEEP-SEA RES (A OCEANOGR RES PAP)* 41, no. 4 (1994): 695-719.

2nd in situ wcs result --
Cape Verde Plateau, E. tropical Atlantic
well above CSH

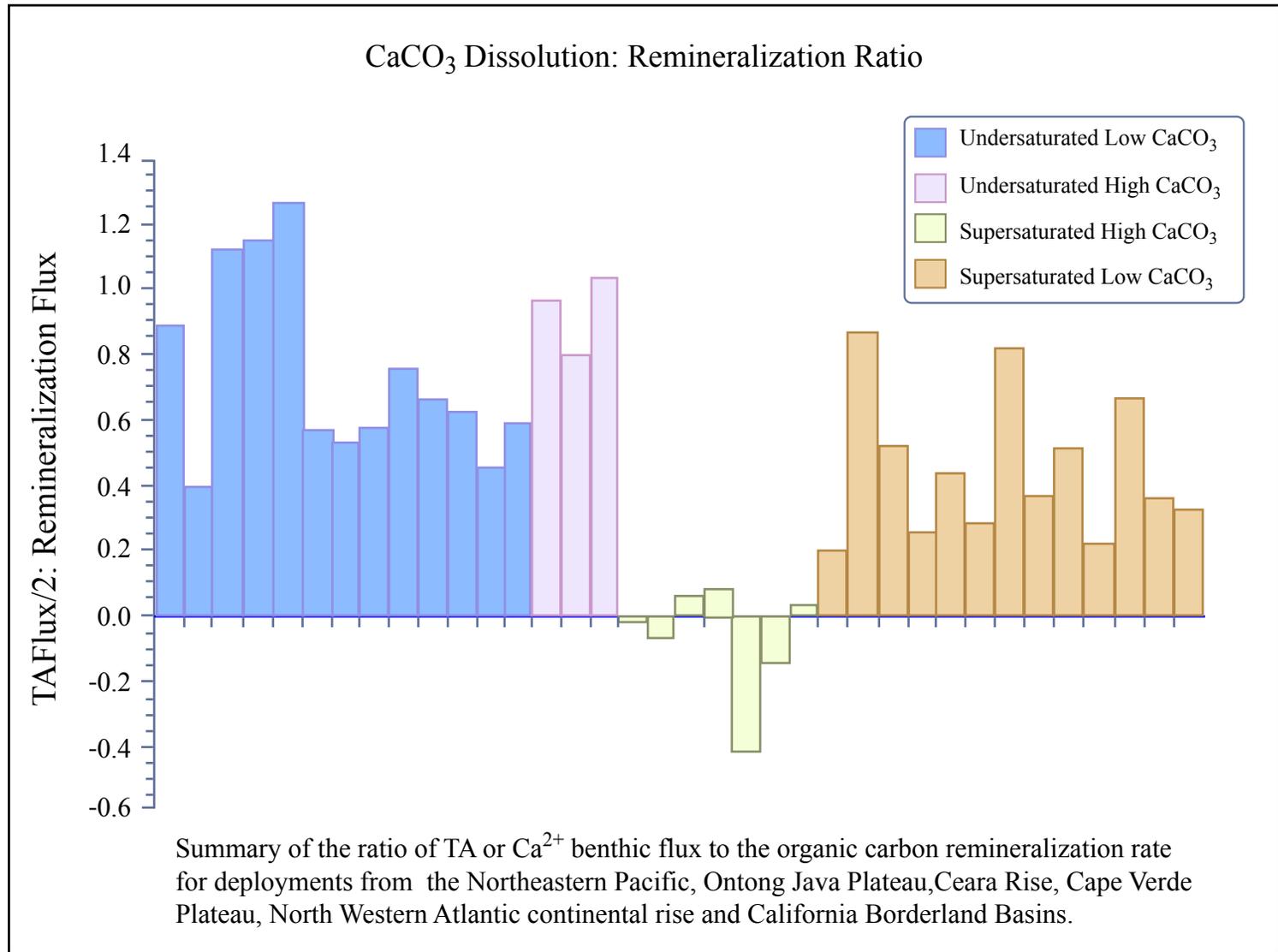
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Lines = fits of model to data to quantify dissolution rate

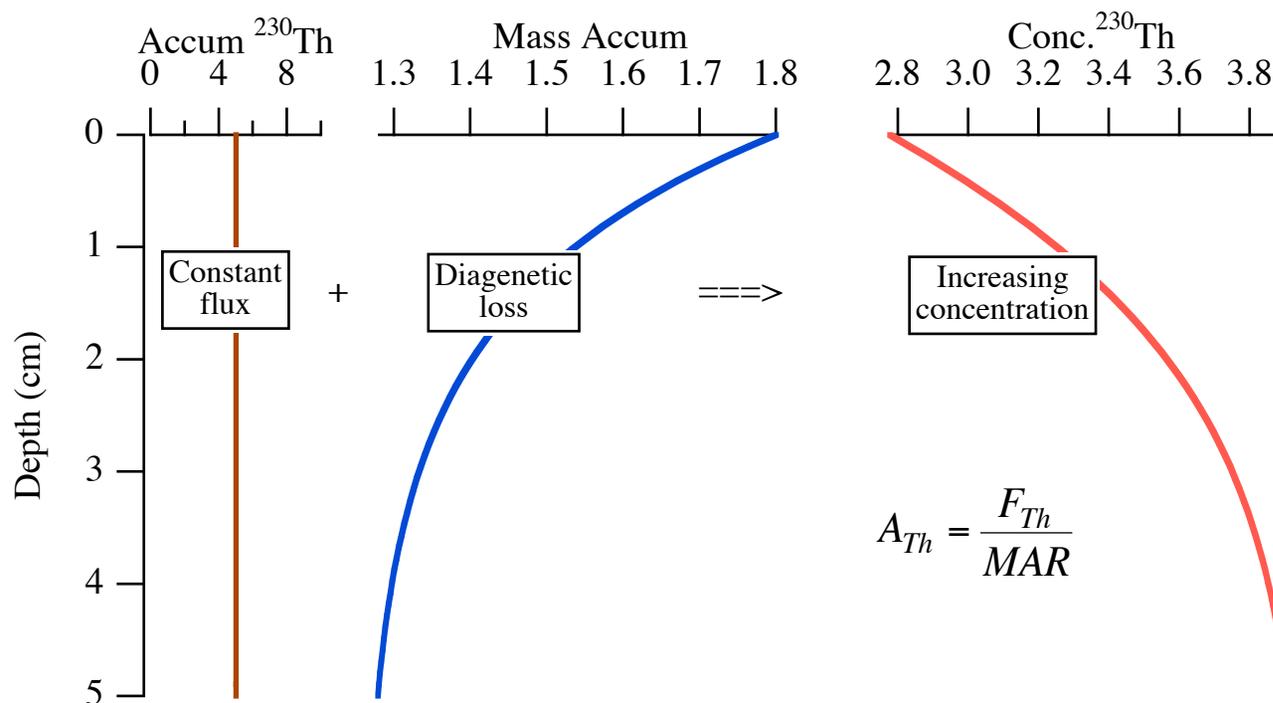
Counter evidence?

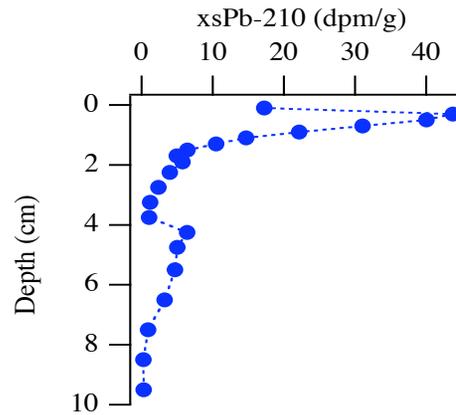
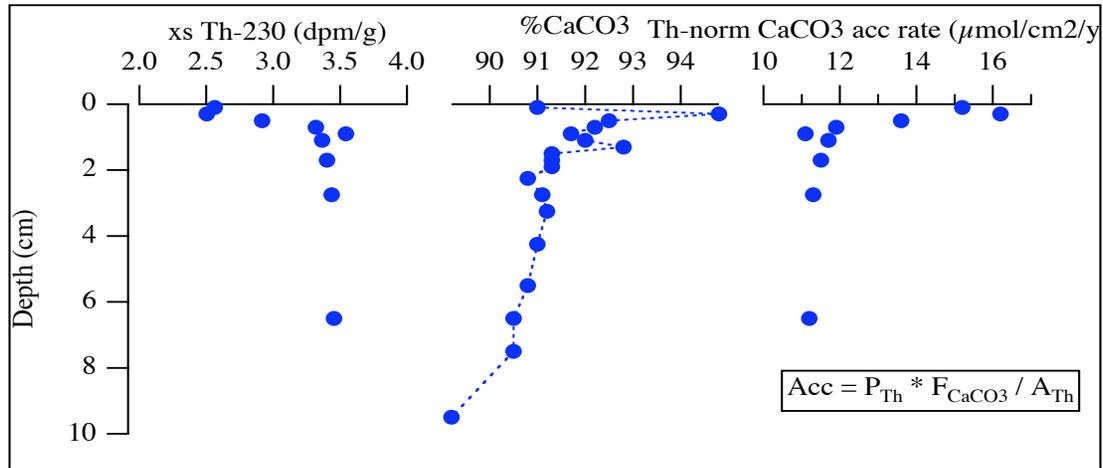
In situ benthic flux chambers

Jahnke & Jahnke, 2004, GCA 68, 47-59



One more approach -- ²³⁰Th activity changes near swi Martin, 2004





Results:
 Depth = 1614m
 BW $\Delta\text{CO}_3 = + 11 \mu\text{mol}/\text{kg}$

Conclusions -- CaCO₃

- 1) Dissolution is driven by undersaturation -- calcite is the most important CaCO₃ mineral in the deep sea.
- 2) Calcite solubility + biogeochemical cycles ==> the degree of saturation of seawater with respect to calcite decreases with increasing depth AND decreases going from the deep Atlantic to the deep Pacific
- 3) Calcite solubility -- that is, its preservation efficiency -- drives the major features of the oceanic calcite distribution
- 4) BUT : oxic metabolism can drive dissolution of calcite in sediments lying above the calcite saturation horizon. This “metabolic dissolution” may play an important role in the marine carbonate cycle -- but its occurrence in high %CaCO₃ sediments is debated.