

## 12.109 Lecture Notes

October 6, 2005

### Petrology

#### Pyroxene thermobarometry

Peridotites contain two pyroxenes: one is opx, orthopyroxene, and the other is a high Ca cpx, clinopyroxene. These two minerals can be used to estimate T + P of the last residence of the nodule. This is thermobarometry!

How do you estimate T + P?

Lindsley & Dixon (graph in handout, En-Di series)

2 coexisting phases—NOT a solvus

Mg-Ca in coexisting pyroxene phases to set T, not very P sensitive.

Also use garnets – pyrope  $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$

Pyx not pure Ca-Mg phases

As P increases, increase solubility of CaTs ( $\text{CaAl}(\text{Si},\text{Al})\text{O}_6$ ) component

In opx, Mg-Tschermakite

Two components in opx add to pyrope:



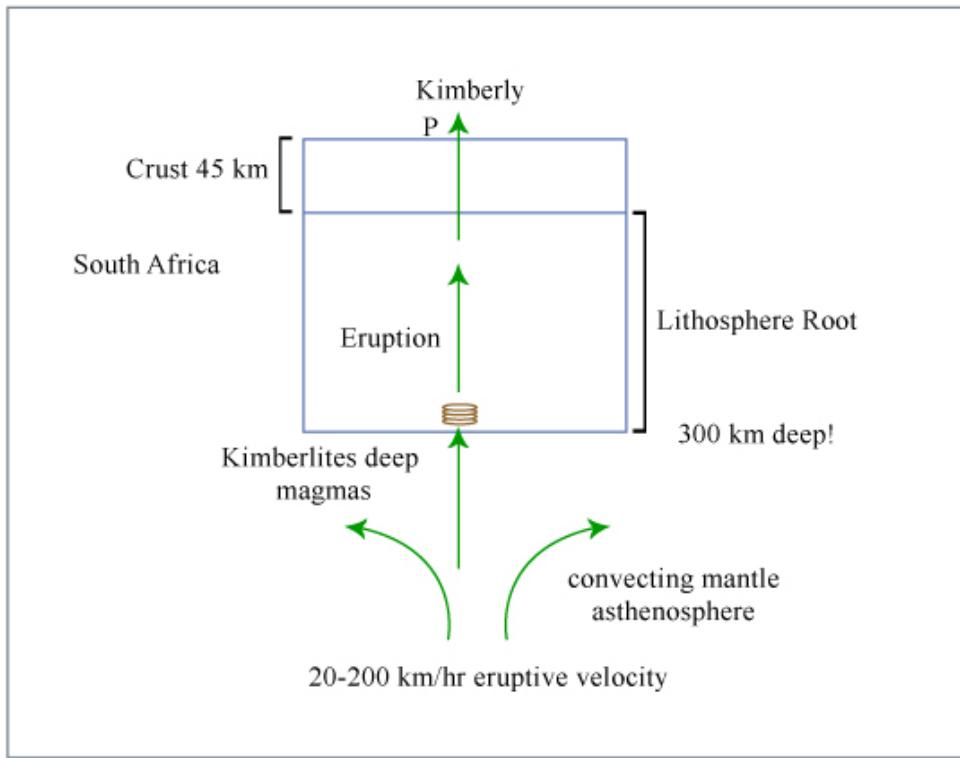
$$k_{eq} = \frac{a_{\text{garnet(pyrope)}}}{(a_{\text{MgAl}_2\text{SiO}_6})(a_{\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}})}$$

$$\ln k_{eq} = \frac{-\Delta G}{RT} = -\frac{\Delta H}{RT} + \frac{T\Delta S}{R} + \frac{(P-1)\Delta V}{RT}$$

This has been experimentally calibrated and is both P + T sensitive.

So, we have two equilibria we can use as thermobarometers in peridotites, Ca-Mg in pyx and garnet-opx Al solubility.

Adiabat –  $\left.\frac{\partial T}{\partial P}\right|_S$  change in P, T for a substance that cannot exchange S (entropy) with the surroundings (an “adiabatic” system).



Kimberlites erupt at 20-200 km/hr velocity

According to Tom Jordan, 1970s Scientific American article, the Kaapvaal craton is >3 billion years old.

How do you keep cratons stable? Buoyancy

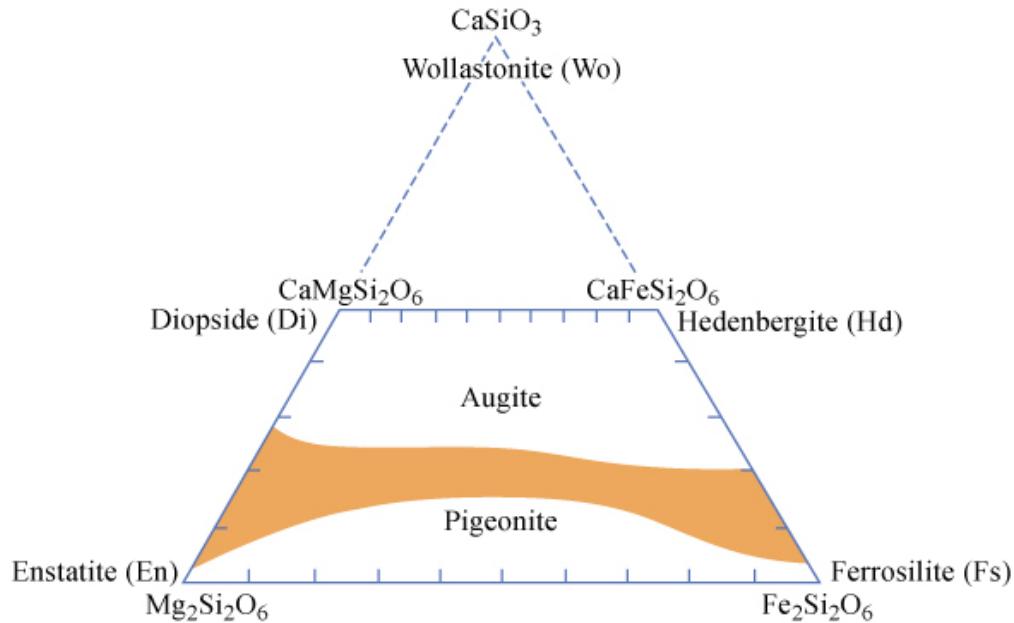
The less dense cratonic material floats on the mantle.

Two types of buoyancy: 1. compositional (Mg lighter than Fe), 2. thermal (hot lighter than cold)

How do you make cratons and keep things cold? Tim says you use a subduction zone. Water subducted with oceanic crust, lowers melting point of the rocks. Wet melting in the mantle wedge can be cooler than would be expected with dry melting. [picture of mantle wedge and volcanoes]

In handout, studies of transition between garnet and spinel peridotite, also studies of pyroxenes.

Pyroxene quadrilateral with isotherms



The pyroxene quadrilateral. For natural pyroxenes with compositions represented on the quadrilateral, the shaded area is a miscibility gap as determined from chemical analyses. Compositions of rapidly quenched metastable subcalcic augites fall in the shaded area between augite and pigeonite. Pyroxenes of virtually any composition within the miscibility gap can be synthesized, but they do not appear to exist stably in nature.

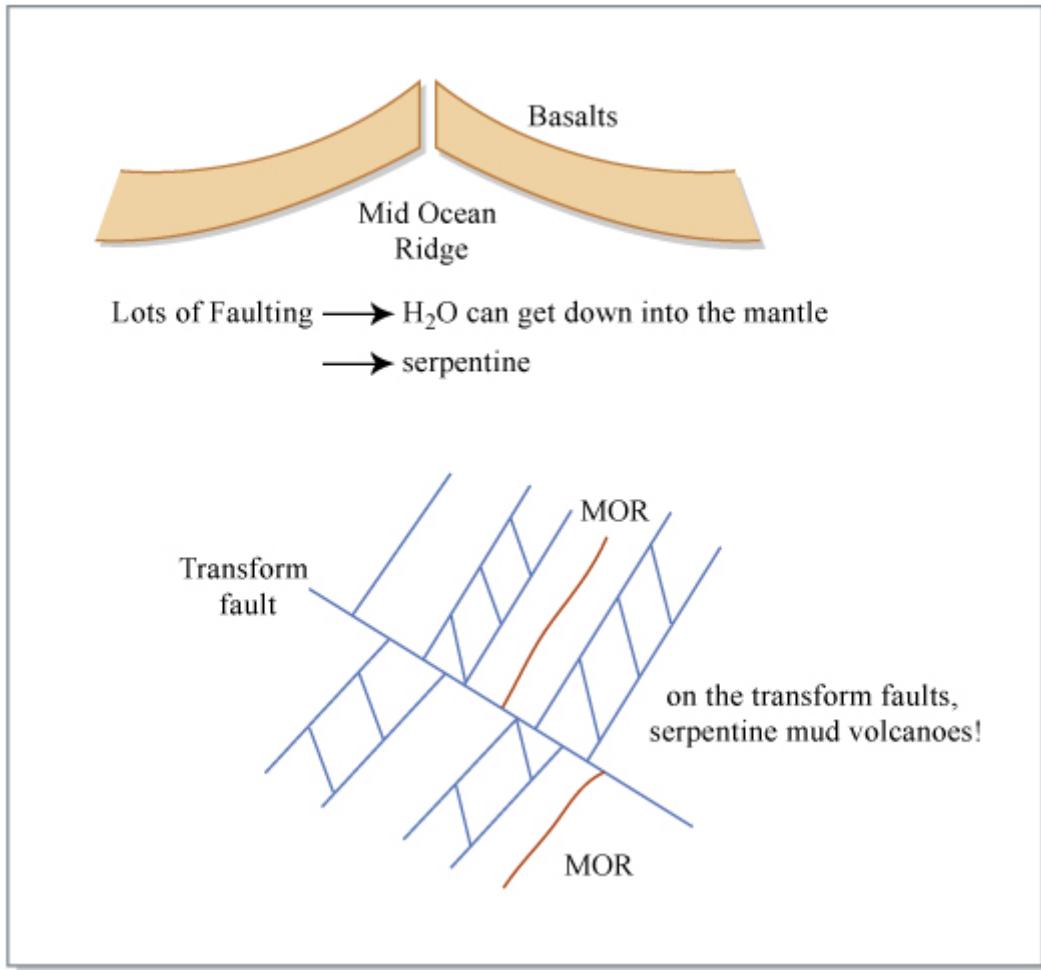
See information on QUILF, DH Lindsey

What happens to peridotite when you metamorphose it in the presence of H<sub>2</sub>O?  
SERPENTINIZATION

Pieces of mantle emplaced tectonically into continental crust are alpine peridotites. Steinmann (Swiss) found association between serpentine peridotite, pillow basalts, and chert in the 1800s, became “Steinmann’s trinity.” Location: Davos, Switzerland. The first explanation for these formations was a hot wet ultramafic magma. Eventually, they realized the serpentine peridotite was hydrated metamorphic rock from the mantle!

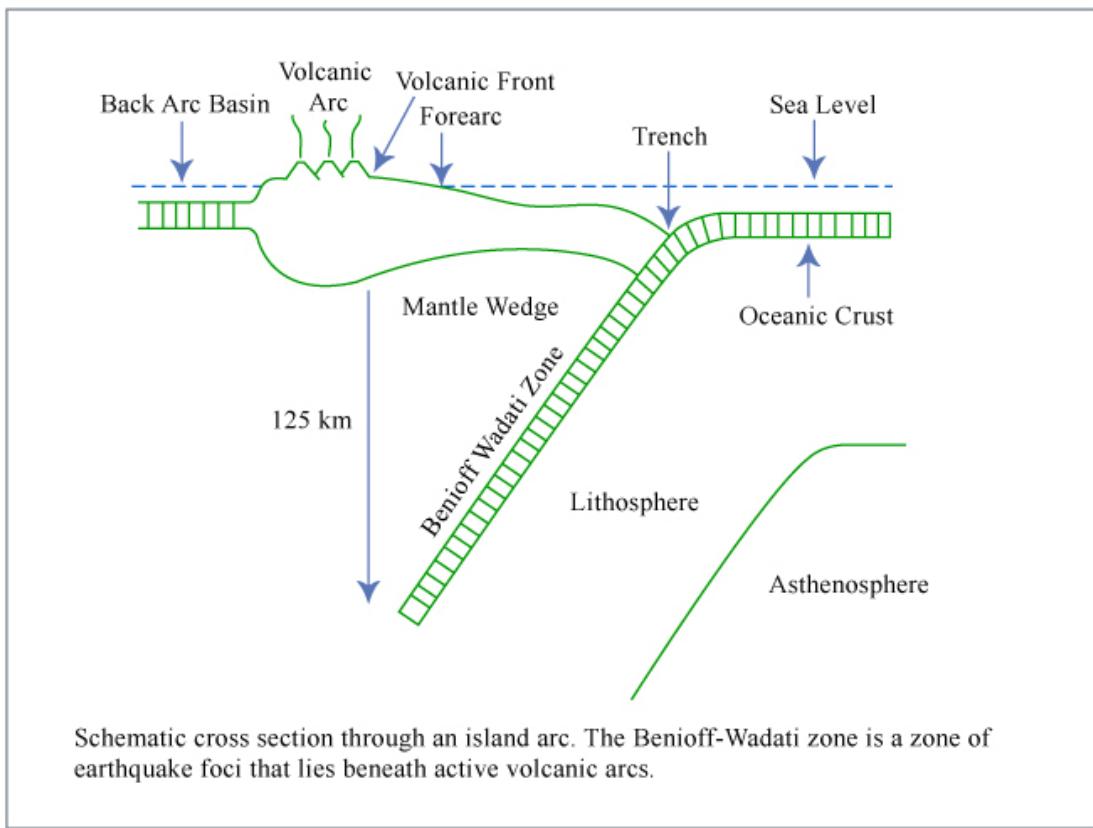
These rocks tell us what happens when you convect mantle that has been near the surface down into the deep earth. The hydrous minerals break down as a function of P + T at ocean spreading centers, and expose peridotite. These are “abyssal peridotites.”

### Abyssal peridotite formation:



Lots of faulting allows H<sub>2</sub>O to get down into the mantle. This results in the production of serpentine at depths up to 20 km!

Serpentine formation at subduction zones:



At depth, serpentine breaks down, releases H<sub>2</sub>O. H<sub>2</sub>O goes up into HOTTER mantle and triggers melting (inverted thermal gradient in mantle wedge)

- 1) dehydration of subducted lithosphere leads to the release of H<sub>2</sub>O  
-rises into overlying mantle wedge, encounters hotter mantle  
-leads to melting
- 2) recycling of H<sub>2</sub>O into the deep earth, phase “A” hydrous silicate

Serpentine is 10% H<sub>2</sub>O, so that's a lot of H<sub>2</sub>O! And serpentine is stable up to high P. Phase “A” (not a mineral yet, still no samples found) is 2% H<sub>2</sub>O.