

November 15, 2005

Bulk composition variations (continued)

Metamorphosed mudstones or shales – pelites

$\text{SiO}_2 - \text{Al}_2\text{O}_3 - \text{K}_2\text{O} - \text{MgO} - \text{FeO}$ system, abundant in tectonically active areas
CaO and Na₂O are not as abundant as these five because they are carried away in solution during surface weathering.

In weathering, plagioclase is the first phase to dissolve from a pelite, through seracic alteration

Calc-silicates

Less abundant, but more attractive are calcareous pelites ($\text{CaCO}_3 + \text{SiO}_2$) or muddy limestones. When metamorphosed, these become calcareous schists (in Europe, called “marl”).

Siliceous Dolomites (siliceous carbonates)

Kaapval craton – 3.1 Billion years ago platform carbonates – stromatolites grow there, might be older stuff but lost since pT.

Only one way to make subduction zones: make pluton-sized granites/rhyoliutes or hydro-alter ocean floor basalts and remelt oceans for a long time
We can't see back that far in the rock record

Meta-basic rocks – metamorphosed mafic volcanics

Ultramafic – serpentinites

Peridotite protoliths

$\text{MgO}-\text{SiO}_2-\text{H}_2\text{O}$

Iron formations – sedimentary rocks, $\text{SiO}_2-\text{FeO}-\text{Fe}_2\text{O}_3-\text{H}_2\text{O}$
BIFs

Metamorphic facies + isograds

1. Isograds

Barrow – mapped metamorphic rocks in Scotland, recognized zones of progressive metamorphism – based on pressure of index minerals

Boundaries between zones marked by the appearance of each index mineral

Pelite sequence: chlorite → biotite → garnet → staurolite → kyanite → sillimanite
See in field: Chlorite zone | Biotite zone, etc.

Boundaries marking the appearance or disappearance of minerals between zones were called isograds.

Isograd – “equal grade” – intersection of a surface of equal P&T with the ground topography

Appearance depends on P, T, fluids, and the bulk composition of a rock

Barrovian metamorphism – classic example of middle P-T conditions (usually in orogenic or tectonic setting)

2. Facies

Eskola (1910-20, later than Barrow) – recognized a distinct set of mineral assemblages that were characteristic of a set of P-T conditions – metamorphic facies

This got complicated because of the large number of named facies. Fortunately, people stopped naming new facies. In common use now:

Barrow's zones	Mafic rock facies	Calc-silicate rock facies
Biotite	Greenschist	Talc-phlogopite
Garnet	Epidote-amphibolite	Tremolite-actinolite-epidote-zoisite
Staurolite		
Staurolite-kyanite	Amphibolite	Diopside-grossular-scapolite
Sillimanite		
K-feldspar	Pyroxene granulite	Forsterite

Plate Tectonics + Metamorphism

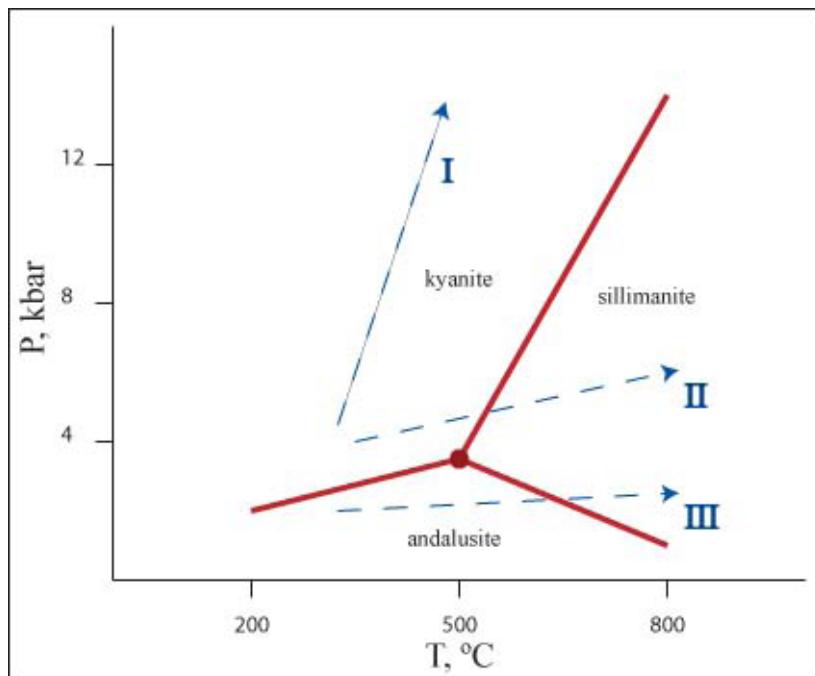
Miyashiro – worked in Japan in 1950s, most famous Ryoke belt – discovered new types of metamorphic rocks

Metamorphic belts were recognized that recorded contrasting P-T conditions

Al_2SiO_5 triple point @ 3.5 kbars, 500°C

1 kbar ≈ 3 km in mantle, 4 km in crust

$10^5 \text{ Pa} = 1 \text{ atm}$



Types

- I. Zeolite → prehnite-pumpellyite → blueschist → eclogite
- II. Barrovian
- III. Low P, high T sequence – andalusite → sillimanite → greenschist → amphibolite

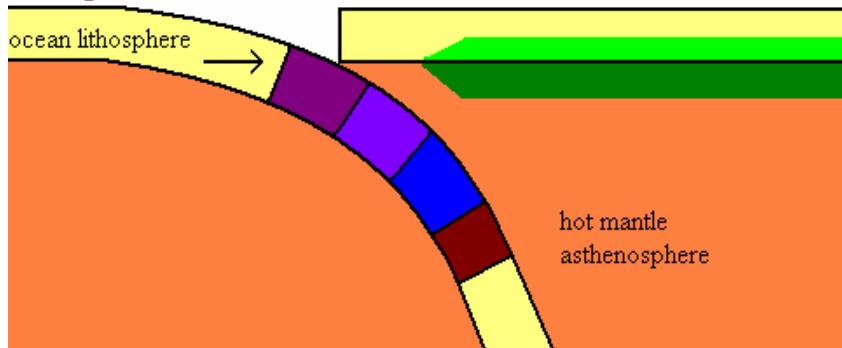
In the 60s when plate tectonics were recognized, it became clear that these zones represented different parts of the subduction environment

I – characteristic of subducted oceanic lithosphere

III – characteristic of the active volcanic arc

Interpretation

- zeolite ■ greenschist
- p.p. ■ amphibolite
- blueschist
- eclogite



3. how do isograds and mineral assemblages record changes in P and T?

system – CaO – Al₂O₃ – SiO₂, 3 components, phase rule $F = c + 2 - \varphi$

F = # degrees of freedom

c = # components

2 = T and P

φ = # phases

so $F = 5 - \varphi$

if $F = 0$ (no degrees of freedom, invariant point), 5 phases coexist

at aluminosilicate triple point, 3 of those phases are the three isomers of Al₂SiO₅